

# 1 Considerations relevant to the stability of granite boulders

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## 9 10 11 12 ABSTRACT:

13 Granite boulders are characteristic geomorphological structures formed in granitic terrains.  
14 Due to their formation process associated with typical spheroidal weathering phenomena,  
15 they tend to show more or less ellipsoidal shapes prone to instability, and they often lie on  
16 small contact surfaces. Analyzing the stability of these boulders is not a straightforward  
17 task. First, these boulders may topple or slide. Additionally, their typically irregular  
18 geometry and uneven contact with the surface where they lie makes the analysis more  
19 complex. The authors have identified some critical issues that are relevant to characterize  
20 these boulders from a rock mechanics point of view, with the aim of estimating the stability  
21 of boulders. In particular, an accurate description of the geometry of the boulder is necessary  
22 to perform accurate toppling calculations. Additionally, the contact area and the features of  
23 the contact plane need to be known in detail. The study is intended to serve as a guideline  
24 to address the stability of these granite boulders in a rigorous way, since standard rock  
25 mechanics approaches (planar failure, toppling stability, standard rock joint strength criteria,  
26 etc...) may not be directly applicable to these particular cases.

27  
28 **Key Words:** Granite boulders, stability, sliding and toppling, geological hazard

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## 32 33 34 **1. Introduction**

35 Large granite boulders are typical geomorphological structures formed in granitic terrains  
36 and are most common in temperate regions that are or have been humid on all continents.  
37 Accordingly, they can be found in the western part of Europe (Galicia in Spain, northern  
38 Portugal, UK), in southern Africa (Zimbabwe, Kenya or Namibia), in southern Asia  
39 including India or Thailand, Brazil, Australia and the USA (e.g. Yosemite or Joshua Tree  
40 National Parks), for example.

42 Due to their formation process, associated to differential weathering, they tend to show more  
43 or less ellipsoidal shapes, which in turn means that they often lie on a small contact surface,  
44 making these boulders prone to instability. Instability of these boulders can put people or  
45 property at risk, meaning this phenomenon should be studied in detail. Furthermore, the  
46 geometry of these blocks tends to be irregular, which contributes to making rigorous  
47 analyses of their stability difficult.

48  
49 The authors have been involved in the stability analysis of some of these structures in the  
50 past and have developed approaches to analyze their stability (Alejano et al., 2010; Pérez-  
51 Rey et al., 2019a), either against sliding or toppling/overturning. At the same time, they  
52 have identified some critical issues that are relevant in order to characterize these boulders  
53 from a rock mechanics perspective to help quantify their stability.

54  
55 Characterization of the contact surfaces of the boulders with the rock where they rest shows  
56 that they cannot typically be considered standard unfilled rough joints, in contrast to most  
57 of the joints usually found in rock masses. Their behavior is best represented by so-called  
58 mismatched joints, where the two contact surface roughness profiles differ.

59  
60 It is also important to note that boulders are usually not regular or symmetric solids.  
61 Accordingly, to compute their stability against overturning, one has to analyze the projection  
62 of the center of gravity of the boulder on the resting plane in relation to the contact base.  
63 Rounded corners also play a role in stability computations.

64  
65 In this paper, relevant features concerning the stability of these boulders against sliding and  
66 toppling based on practical experience will be briefly reviewed and illustrated with the help  
67 of physical models and actual case studies. Based on previous applied studies (Alejano et  
68 al., 2010; Pérez-Rey et al., 2019), some improvements were incorporated and new analyses  
69 area provided to insight granite boulder stability behavior. This paper is ultimately intended  
70 to be of help in addressing the stability of granitic and other type of boulders for rock  
71 mechanics practitioners dealing with this type of problems, since standard rock mechanics  
72 approaches may not be directly applied to this particular type of stability analyses.

## 73 74 75 **2. Granite boulder formation and occurrence**

76 This study addresses instability phenomena associated with irregular stone boulders in  
77 granitic terrains, illustrated in the context of the conditions encountered in the northwest of  
78 the Iberian Peninsula. These phenomena also take place in other regions of the globe where  
79 granitic basements are subjected (now or in the past) to intense weathering. A survey of  
80 granite landforms from a geomorphologic perspective was produced by Twidale (1982),  
81 where the author refers to four main groups of landform types including boulders,  
82 inselbergs, all-slopes topography and plains.

83

84 Disregarding the stable plains, three types of potentially unstable slope environments can  
85 be identified among these landforms, including typically large ellipsoidal individual granite  
86 boulders areas, whose stability is the main focus of this study, mid-slope regions formed by  
87 groups of medium-size granitic boulders in decomposed granite matrix (which can produce  
88 rockfall phenomena; Pérez-Rey et al., 2019b) and, finally, mild slopes formed by highly or  
89 completely decomposed granite (HDG or CDG) (Jiao et al., 2005; Alejano & Carranza-  
90 Torres, 2011; Jiao et al., 2012; Ohtsu et al., 2018) (Figure 1). In all these cases, weathering  
91 of granitic materials plays an important role.

92

93 FIGURE 1

94

95 Boulders are one of the most common and characteristic landforms of granitic terrains and  
96 they originate due to weathering (Durgin, 1977; Fletcher et al., 2006). These structures form  
97 through the mechanism of spheroidal weathering and forward erosion of decomposed  
98 granite. Spheroidal weathering is a physical-chemical process that affects uniform rock  
99 masses with regular joint patterns, typical of granitic rocks, but also in volcanic tuffs or  
100 basaltic rock.

101

102 Weathering is the process by which rock deteriorates until it eventually breaks down to a  
103 soil. This process is highly dependent on climatic influences (Selby, 1993). Often,  
104 [weathering works from free surfaces](#) where chemicals in water attack the parent rock  
105 (Figures 2 and 3). Eventually it may leave a framework or core-stones of more or less fresh  
106 rock separated by weathered zones that can be easily eroded (Ollier, 1975; Hack, 2009; Md  
107 Dan et al., 2016). Often, joint sets found in rock masses are orthogonal; two sets occur  
108 perpendicular to one another and perpendicular to some planar fabric such as bedding,  
109 foliation or flow banding in an igneous pluton (Taboada et al., 2005; Hencher, 2015).

110

111 Linton (1955) theorized a two-stage process of formation (Figure 2). The first stage involves  
112 deep penetration of weathering along joint surfaces, which produces a thick saprolite or  
113 completely decomposed granite (CDG) mantle interspersed with non-weathered core-  
114 stones. The second stage is brought on by exhumation either by tectonic uplift of these  
115 boulders or lowering of base level accompanied by erosive processes.

116

117 FIGURE 2

118

119 FIGURE 3

120

121 In the case of granites, rock erosion tends to produce boulder fields or the so-called tors or  
122 inselbergs, which are residual rock masses that display as isolated piles of boulders (Twidale  
123 & Vidal Romani, 2005). This is consistent with the formation process suggested by Linton  
124 (1955).

125

126 Following weathering, granular saprolite or completely decomposed granite (CDG) is  
127 quickly eroded or removed by wind and water leaving behind the rounded core-stones or

128 boulders. An example of one of these processes is illustrated in Figure 3, where a structure  
129 of some granitic boulders eventually remains after weathering of a jointed rock granitic  
130 mass.

131

132 If these erosive processes persist until they reach the entire rock mass, a phenomenon typical  
133 of granite plains, the granite becomes completely decomposed, behaving like a soil material  
134 (HDG or CDG); this is also known in the NW of the Iberian Peninsula as “jabre” (Alejano  
135 & Carranza-Torres, 2011), but has other local names according to geography (GEO, 1988;  
136 Onitsuka et al., 1985). One defining aspect of CDGs in the context of geotechnical  
137 engineering is their heterogeneous spatial distribution and natural variability (Figure 1.c).

138

139 Since the original joints of a granitic mass are not necessarily orthogonal (though they tend  
140 to be), granite boulders occur in different shapes and sizes, from almost perfect spheres to  
141 ellipsoidal bodies, but also slender or irregular slabs. As previously noted, these boulders  
142 are common all over the world, but particularly in temperate humid regions (Figure 4).

143

144

FIGURE 4

145

146

### 147 **3. Granite boulder instability mechanisms**

148 Granite boulders are effectively rock blocks with complex geometry lying on approximately  
149 planar resting surfaces. In the field of rock slope engineering, the stability of a rock block  
150 lying on an inclined plane should be studied against the two types of instability mechanisms  
151 hypothetically observable in these cases, namely, sliding and toppling/overturning  
152 (Sagaseta, 1986). If we consider that a granite boulder is a block lying on such an inclined  
153 plane, both types of mechanisms should be considered (Figure 5) in order to quantify its  
154 stability level.

155

156

FIGURE 5

157

158 Traditional rock slope engineering studies (Hoek & Bray, 1974; Goodman & Bray, 1976;  
159 Wyllie & Mah, 2004) developed methodologies for analyzing the stability of rock blocks  
160 with simple shapes delimited by pre-existing rock joints or discontinuities, such as rock  
161 slabs, prisms, wedges or columns. These potentially unstable blocks tend to form when  
162 excavating man-made slopes or rock cuts. Stability against sliding or toppling of complex  
163 geometry boulders or blocks is not a straightforward task since adapting calculations to the  
164 observed geometries can be difficult. Additionally, contact zones between boulders and  
165 resting planes do not tend to behave like standard unfilled rough rock joints (further  
166 explained in section 4.6), so traditional rock joint strength approaches may not be  
167 appropriate.

168

169 The stability against sliding, in case the contact plane is cohesionless (as is normally the  
170 case), is controlled by the plane dip ( $\alpha$ ) and the friction angle ( $\phi$ ) of the contact between the  
171 boulder and the basal plane according to Eq. 1:

$$172 \quad FOS_{sliding} = \frac{\tan \phi}{\tan \alpha} \quad (1)$$

173  
174 In the case of rough unfilled rock joints, the Barton-Bandis approach (Barton & Choubey,  
175 1977; Barton & Bandis, 1982) can be used to compute the friction of the contact, but the  
176 authors' experience is that the boulder-rock contact does not behave like this type of joint  
177 (Alejano et al., 2012).

178  
179 The basic equation controlling the stability against toppling of a rigid block is presented in  
180 Eq. 2 and can be used to estimate the factor of safety, and accordingly, the stability of a  
181 block against toppling.

$$182 \quad FOS_{toppling} = \frac{\sum M_{stabilising}}{\sum M_{overturning}} \quad (2)$$

183  
184 This simple equation just considers the ratio of the stabilizing and overturning moments in  
185 relation to the corresponding rotation axis. In the simplest case, where the only driving force  
186 is the weight of the specimen, the factor of safety against toppling ( $FOS_{toppling}$ ) can be  
187 computed according to the forces acting along x and y-axes, in relation to a rotation axis  
188 located at the lower corner of the block in the direction of tilting (Figure 5.a).

189  
190 The forces involved in the analysis of each specimen result only from its own weight (and  
191 other external forces) and are typically applied at the center of gravity of the specimens. For  
192 slabs with simple geometry, the rotating axis is easily identified as the lower external corner,  
193 but when the corners of the slab or boulder are rounded or the boulder has complex  
194 geometry, the rotation axis should be chosen with care. Indeed, for boulders with an  
195 irregular base, the rotation point varies according to the projection of the center of gravity  
196 on this base.

197  
198 The most likely failure mechanism (sliding or toppling) will be the one theoretically  
199 occurring at a lower tilting angle. Typically, slender boulders will be more prone to topple,  
200 whereas rounded blocks tend to slide, if the basal plane dip is larger than the apparent  
201 friction angle of the contact (Figure 6).

202  
203 **FIGURE 6**

204  
205 With the aim of studying the stability of rock elements, it is possible to carry out simple tilt  
206 tests under controlled environmental conditions and constant lifting velocities to estimate  
207 analytically predicted angles in the laboratory (Alejano et al., 2015 & 2018).

208

209

## 210 **4. Relevant issues affecting stability calculations**

211 In this section, the authors address a number of relevant issues to be considered when  
212 estimating boulder stability, adapting classic rock mechanics approaches to the types of  
213 cases under scrutiny.

214

215

### 216 4.1. Detailed boulder geometry

217

218 One of the reasons that has made it difficult in the past, if not impossible, to compute boulder  
219 stability was the unavailability of a detailed knowledge of the geometry of the boulders.

220

221 Advanced surveying techniques such as Terrestrial Laser Scanning (TLS) and close range  
222 or drone (remotely piloted aircraft systems, RPAS) photogrammetry permits a very detailed  
223 record of the land geometry in the form of 3D point clouds (3DPC) (Armesto et al., 2009;  
224 Ferrero et al, 2008; Riquelme et al., 2014).

225

226 Our experience is that it is better to apply TLS (generally more accurate) and RPAS  
227 photogrammetry together, since due to the shape and size of the boulders, their upper part  
228 is usually hidden from terrestrial views and their lower parts and contact zones are shadowed  
229 in top views typical of RPAS photogrammetry. By combining information recovered from  
230 both in-situ non-contact surveys, an accurate 3DPC can be obtained. This 3DPC can be  
231 processed using software like MeshLab or CloudCompare (Girardeau-Montaut, 2018)

232

233 Figure 7.a illustrates the ‘Pena do Equilibrio’ boulder studied by Pérez-Rey et al. (2019a)  
234 and Figure 7.b shows the 3D point cloud derived from TLS and UAS imagery obtained for  
235 this boulder, where also its center of gravity and the relatively very small contact area are  
236 depicted. This information is critical for further sufficiently accurate calculations. Figure  
237 7.c. represents the top view of the boulder together with the contact zone and center of  
238 gravity (*cog*) projections (also enlarged). This projection is needed to compute stability  
239 against toppling of this block, as explained below.

240

241

### 241 FIGURE 7

242

243

### 244 4.2. Stability against toppling

245

246 Using detailed geometry data (e.g. a 3DPC) including contact area, the stability of the  
247 boulder against toppling can be reliably computed.  $FoS_{\text{toppling}}$  is computed according to Eq.  
248 2. Application of this formula to an idealized slab geometry is straightforward as shown in  
249 Figure 8, derived from the seminal Goodman & Bray (1976) approach.

250

251 The location of the rotation axis is easily identified, as it is located on the lower outer corner  
252 of the rectangular-shaped slab. If the projection of the center of gravity (*cog*) falls within  
253 the base of the block, it will not topple, whereas when it falls out of this base, the block will  
254 topple.

255

256 To analyze the stability of a slab with rounded or eroded corners (Figure 9), the situation  
257 becomes slightly more complicated. However, if the rounding corner radius ( $r$ ) is known,  
258 computations can be performed according to Alejano et al. (2015) and the formulation  
259 presented in Figure 9. Rounding of the corners contributes to making boulders more prone  
260 to toppling.

261

FIGURE 8

262

263

FIGURE 9

264

265

266 To illustrate this effect of rounding corners, a slab-like solid with side lengths of 3, 4 and  
267 5 cm and rounded corners with 1 cm curvature radius have been printed in a heavy plastic  
268 by means of a 3D printer (Figure 10). This small block has been tilted in the 12 possible  
269 positions (4 corresponding to the slenderness ratio 5/3, 4 for 4/3 and 4 for 5/4). The average  
270 tilt angles obtained for every three groups of four tests, together with the theoretical value  
271 obtained applying the equations of Figure 9 are compiled in Table 1, showing a very good  
272 agreement, confirming the validity of the round corner toppling stability computation  
273 approach.

274

FIGURE 10

275

276

TABLE 1

277

278

279 Remark that for the case of actual boulders the so called curvature radius may vary along  
280 an edge and in different edges of the boulder, so the selection of a representative value of this  
281 parameter, tending to diminish in the middle of the edge, may slightly affect the stability  
282 computations (Alejano et al., 2015).

283

284 To further illustrate the influence of geometry on toppling stability, a number of increasingly  
285 complex geometrical elements are illustrated on the top row of Figure 11 with the aim of  
286 analyzing their toppling behavior.

287

288 Figure 11.a is a simple rectangular prism with square base. If we position it on a flat surface,  
289 its center of gravity (*cog*) will project right on the center of its square base. If we  
290 continuously tilt the surface where it stands, at a particular tilt angle  $\alpha$ , the projection of its  
291 *cog* will come out of its base (lower row of Figure 11) so the prism will topple. The tilting  
292 angle  $\alpha$  for toppling can be computed based on the formulation illustrated in Figure 8.

293

FIGURE 11

294

295

296 To see the role of a more complex geometry, Figure 11.b shows a similar prism with a cube  
297 stuck to its upper back face. Note that when placed on a flat surface, its *cog* will not project  
298 on the center of the square base but somewhat backwards, due to the effect of the additional  
299 weight of the stuck cube. In this way, when tilting the plane on which this element stands,  
300 it will topple at a higher angle than the previous figure because the *cog* will project  
301 somewhat backwards. In this way the toppling angle in this case,  $\beta$ , will be steeper than the  
302 one observed for the first prismatic element ( $\alpha$ ).

303

304 Element *c* in Figure 11 is like element *b*, but the added cube is now stacked on the upper  
305 left side face. Its *cog* will be at the same height as for element *b* (since it is the same element),  
306 but its projection on a horizontal plane will be moved leftwards in relation to element *a*  
307 (Figure 11.c, second row). When tilting the platform where element *c* rests, it will topple at  
308 a less steep angle than  $\alpha$ , because its *cog* is located higher than in case “*a*”, so its projection  
309 will fall outside its base at a less steep angle  $\gamma$ , which will be also less steep than  $\beta$ .

310

311 Element *d* in Figure 11 is a rectangular prism with square cross-section that has cubes  
312 attached in the upper part of its lateral backward and leftward faces. In this case, the *cog* of  
313 the element will be even higher than for elements *b* and *c* and the *cog* projection on its base  
314 will be slightly moved backwards and a little bit leftwards in relation to the case of element  
315 *a* in Figure 11. This will be clearly less stable than *b* (since the side-stuck cube moves the  
316 *cog* upwards), but more stable than *c* (since the back-stuck cube will increase its stability by  
317 moving the projection of its *cog* backwards).

318

319 One can then show that  $\beta > \delta > \alpha > \gamma$ . Based on this type of reasoning or equivalent  
320 computations, it is possible to develop an understanding of general toppling mechanics  
321 using the contact area and projection of the *cog*. This approach can be applied to better  
322 understand stability of granite boulders.

323

324 Note that when no external forces (e.g. water or seismic forces) are applied, the stability of  
325 these elements basically depends on the location of the vertical projection of the *cog* of the  
326 element on the contact base. When the projection of the *cog* is inside the contact base, the  
327 element is stable against toppling; when this projection falls out the contact base, the  
328 element will topple. External forces typically associated with seismic movements or water  
329 pressure can destabilize otherwise blocks.

330

331 Considering this, a factor of safety against toppling can be computed by relating the angle  
332 between the vertical line and the normal to the basal plane (denoted as  $\alpha$  in Figures 8, 9 and  
333 12) and the angle formed by the normal to the basal plane and the line connecting the center  
334 of gravity and the rotation axis. Note that the rotation axis is the point where the projection  
335 of the *cog* will first come out of the contact area when tilting the basal plane (shown as  $\beta'$   
336 in Figures 9 and 12).

337



## FIGURE 12

This concept is illustrated for the ‘Pena do Equilibrio’ boulder studied by the authors (Pérez-Rey et al., 2019a) in the enlarged view of Figure 7.d and in Figure 12. In both these figures, the projection of the *cog* normal to the basal plane (*cog-p*, in brown in the figures) falls out of the contact surface. This means that if the boulder were placed on a horizontal base, it would have toppled backwards. The vertical projection of the *cog* on the contact base (*cog-v*, in green color in Figures 7.c and 12) falls within this contact area, which explains the present stability of the boulder against toppling.

Moreover, the factor of safety of the boulder against toppling, denoted as  $FO_{Stoppiling}$ , can be computed as the relation of the tangents of the angles indicated above (shown as  $\beta'$  and  $\alpha$ , respectively, in Figures 9 and 12). This approach can be extended to the case where external forces such as water pressure, ice-jacking or a seismic force are applied to the boulder.

It should be noted, however, that such potential stability estimates depend on a very accurate description of the geometry of the boulder, including its contact area. Also, knowledge of the geometry of the basal plane (dip, dip direction and planarity) and the contact zone between the boulder and basal plane area and external forces acting is needed to assess stability. When the contact zone has a simple geometry (Figures 8, 9 and 12), identifying the rotation axis is simple. However, for less regular contacts, as are usually found in nature, the situation is more complex, as illustrated in section 4.4.

### 4.3. Positioning of the rotation axis for toppling estimates

Often, the location of the rotation axis enabling stability calculations of boulder toppling may not be known *a priori*. This is the case of the boulder illustrated on Figures 7 and 12, or any other element that does not have an edge parallel to the strike of the basal plane.

To illustrate how to identify the rotation axis, a simple tilt test is performed with an element consisting of two pieces: a cylindrical rock specimen (with height twice its diameter) and a steel disk with the same diameter but much denser than the rock, which is positioned on the rock cylinder and moved a half radius leftwards as illustrated in Figure 13.a. In this way, the projection of the *cog* of the element moves towards the left and it does not project on the center of the rock cylinder base.

When this element is positioned on a tilting platform and increasingly tilted (Figure 13.b), the projection of the *cog* moves along the dip direction of the tilting table until it projects on the perimeter of the rock cylinder base. At this point, the whole setup will topple. This toppling will not occur in the dip direction, but in a direction forming an angle  $\beta$  with the dip direction as shown on Figure 13.b and demonstrated by Pérez-Rey et al (2019a).

382 This means that the toppling will occur in a direction marked by the point where the *cog*  
383 first projects outside the element base as shown in the lower left corner photograph in Figure  
384 13.b. Accordingly, if the boulder illustrated on Figure 12 topples, it will not do so in the dip  
385 direction, but in the direction marked by this intersecting point.

386

387 Additionally, the rotation axis necessary to compute stability against toppling will have to  
388 be identified based on this intersection. Pérez-Rey et al (2019a) illustrates how this is  
389 performed in greater detail.

390

391

## FIGURE 13

392

### 4.4. Contact zone geometry

393

394  
395 In all the reasoning above, we have assumed a fully planar contact surface between the  
396 element or boulder under scrutiny and the so-called basal plane. Even if this seems to be a  
397 reasonable assumption, it may not always hold true, which can may positively or negatively  
398 affect the boulder stability. Specifically, a concave contact will improve stability against  
399 toppling, and a convex contact will be detrimental. However, the non-planarity of this  
400 contact is difficult to characterize in practice.

401

402 In theory, at least for specific tested geometries, the convex or concave contact geometries  
403 could be incorporated into stability computations against toppling. For illustrative purposes,  
404 the authors have checked this formulation for the elements shown in Figure 14 and compare  
405 analytical results against actual tilt tests results, based on the force distribution and pictures  
406 of the two tested elements in Figure 14.

407

408 The authors tentatively suggest that a convex contact can be associated with traditional  
409 rocking stones, also known as logan stones or logans. These are large stones that are so  
410 finely balanced that the application of just a small force causes them to rock. These rocks  
411 associated with popular traditions appear in the Atlantic European façade and elsewhere as  
412 pictured in Figure. 4.b.

413

414

## FIGURE 14

415

416 Not only the concavity of convexity of the base contact can affect stability, but also  
417 roughness may play a role. Indeed, typical roughness of rock joints could affect the critical  
418 value of the tilting angle, according to the wavelength and asperity amplitude, and  
419 particularly the asperity height in the potential rotation axis. Variations of a few degrees in  
420 the critical tilting angles and variations on the order of 10% of the  $FO_{S_{toppling}}$  can be observed  
421 based on the roughness of the basal plane.

422

423 To illustrate this phenomenon from a theoretical point of view, a granitic composite  
424 specimen is considered with a regular rough base (Figure 15). It has a height of 99.17 mm  
425 in the valleys and 93.64 mm in the peaks with a width of 46.36 mm. The roughness presents

426 an inclination of 20° (equivalent JRC=20) at a wavelength of 30.92 mm and an amplitude  
427 of 8.53 mm. When tilting this specimen in both possible directions, it topples at 20.9° when  
428 the toppling corner coincides with a valley, and at 23.6° when the rotation axis coincides  
429 with a peak. The corresponding theoretical estimates are 24.5° and 27.1°; once corrected  
430 considering a round corner with radius of 7.5% the width of the samples (a value selected to  
431 match the experimental values through back analysis), these values change to 21.1° and 23.5°,  
432 obviously quite close to those that were experimentally observed.

433

434 This simple study shows that the topology of the granite boulder base, if rough, could  
435 influence the stability condition of the boulder in relation to toppling. For the simple case  
436 introduced here, the variations introduced to the critical toppling angle are about 3°; this  
437 effect will be less significant for slenderer samples, and typically more significant when the  
438 contact zone size is small.

439

#### 440 FIGURE 15

441

#### 442 4.5. 3D printing of boulders and physical modelling

443

444 Overturning of a constant density rigid body only depends on its geometry and its position  
445 on a basal plane. If the body is formed by materials with varying densities, it will also depend  
446 on the location of the *cog*. Therefore, if one can reproduce the geometry of a boulder with a  
447 different material, for instance any plastic as those used by 3D printers, it is possible to carry  
448 out physical models including tilt tests to analyze stability against toppling of the boulder.  
449 In the case that, as recommended, we have a detailed and accurate point cloud (3DPC) of  
450 the boulder available, it is now feasible to print a scaled 3D version of the boulder under  
451 consideration.

452

453 The authors have created a roughly 1:50 scaled plastic version of the ‘Pena do Equilibrio’  
454 boulder illustrated in Figure 7. Figure 16 shows the upper and lower plan views of the  
455 boulder, including the contact area and the boulder in the process of subjecting it to a tilt  
456 test. The polylactide (PLA) plastic replica of the boulder presented a plastic pattern inside,  
457 but we consider the assumption of uniform density to be reasonable. It is therefore an  
458 appropriate geometrical copy of the actual boulder for use in physical testing.

459

#### 460 FIGURE 16

461

462 Tilt tests with this boulder, adequately positioned on a polystyrene surface and with sand  
463 paper in the contact area, provided tilt test toppling angles in the range of 30 to 31°, one  
464 degree (on average) less than the critical tilting angle estimated for the actual boulder based  
465 on analytical calculations.

466

467 As an alternative to the physical modeling demonstrated here, 3D point clouds may serve  
468 as a basis to create a grid able to simulate the behavior of the boulder by means of the  
469 Discrete Element Method or any other suitable numerical modelling technique.

470

471

#### 472 4.6. Contact strength

473

474 The Barton & Bandis (1982) formulation has been widely used as a suitable approach to  
475 estimate the shear strength of rough unfilled rock joints. However, in the process of boulder  
476 formation associated with spheroidal weathering, the original joint typically erodes  
477 differently on its sides which tends to produce convex, and not planar, profiles. Accordingly,  
478 the final contact between the boulder and the basal plane behaves more like a so-called  
479 mismatched contact or joint (Figure 17). In this case, the sides of the contact do not match,  
480 as each side presents a different roughness pattern (*JRC*) and shear behavior tends to depend  
481 more on the contact area than on the *JRC*.

482

483 For illustrative purposes, Figure 18 shows two natural rough joints (left hand side) and two  
484 block contacts (right hand side). Natural joints formed in a rock mass tend to present equal  
485 geometrical patterns in both sides, so they are matched joints and behave accordingly  
486 (Barton-Bandis); conversely, block contacts present different geometric patterns in both  
487 sides, so they are mismatched joints and behave as such.

488

489 FIGURE 17

490

491 FIGURE 18

492

493 Zhao (1997a, b) studied strength behavior of mismatched contacts and proposed a new  
494 version of Barton's formula denoted the *JRC-JMC* shear strength model. This criterion  
495 accounts for an additional influence of the so-called joint matching coefficient (*JMC*), a  
496 parameter to be estimated based on the matching of the two joint sides:

$$497 \tau_{JRC-JMC} = \sigma_n \tan \left[ \phi_r + JMC \cdot JRC_n \log_{10} \frac{JCS_n}{\sigma_n} \right] \quad (3)$$

498

499 Some studies have shown that *JMC* depends on the level of contact of the surfaces. Based  
500 on an estimation of potential contact of such a surface in case of granite blocks or granite  
501 boulders, previous experience developed by the authors with tilt tests carried out on large  
502 scale physical models (Alejano et al., 2012) and on the recommendation by Zhao (1997b),  
503 we tentatively suggest *JMC* = 0.3 as an initial estimate for this type of problem. The authors  
504 think that more detailed studies on the behavior of these mismatched joints will be necessary  
505 to better understand and bracket the shear strength behavior of these contacts.

506

507

#### 508 4.7. External forces

509

510 A detailed analysis of the external forces, including water pressures, ice-jacking and  
511 earthquakes potentially acting on the studied rock structures should be accounted for when

512 analyzing the stability of boulders for different temporal horizons (Christianson et al., 1995;  
513 Alejano et al., 2010).

514

515 Most often, water pressure associated high rainfall periods can be considered, but seismic  
516 forces and ice-jacking are also typical external forces. When such forces are applied to  
517 boulders or rock elements at particular moments, they may cause the ultimate instability of  
518 the element under scrutiny (Wyllie and Mah, 2004; Alejano et al., 2013).

519

520 Meteorological records informing on peak rainfall and lowest levels of temperatures and  
521 freezing periods can help to provide realistic assumptions regarding the magnitude and level  
522 of external forces associated to water on boulders. Additionally, seismic safety acts and  
523 earthquake damage mitigation policies could be of help in order to quantify the role of  
524 earthquakes on the stability of this type of structures.

525

526 Recent studies have put forward other potential external causes of ultimate instability (Vann  
527 et al., 2019). These include dynamic loads caused by construction equipment or seismic  
528 shaking, loss of downslope support, and human activity. All of these and other external  
529 influences should be considered in stability estimates.

530

531

#### 532 4.8. Numerical modeling

533

534 Numerical analysis can also be of help to carry out a realistic assessment on the stability of  
535 boulders with complex geometries. Some authors (Christianson et al., 1995; Shi et al., 1996;  
536 Purvance et al., 2009) have specifically addressed the stability of boulders in regard to  
537 seismic forces in regions particularly prone to earthquake activity.

538

539 The 3D geometry of boulders can be obtained in the form of a 3DPC. This 3DPC can be  
540 typically imported to 3D software to perform stability analysis. For the case of the ‘Pena do  
541 Equilibrio’ boulder, the 3DPC was imported to the finite-element software MIDAS GTS  
542 NX v.2019 (Midas IT, 2019) to generate a 3D tetrahedral mesh with more than 20,000  
543 elements (including boulder and basal rock where it stands). These elements were exported  
544 as individual blocks to the Discrete Element Method-based 3-Dimensional Distinct Element  
545 Code 3DEC v.5.20 (Itasca, 2019) by means of an internal subroutine (Muñiz-Menéndez et  
546 al., 2020).

547

548 For these numerical studies, geomechanical parameters such as the normal and shear  
549 stiffness in the contact plane should be selected with care according to estimative techniques  
550 (Itasca, 2019; Muñiz-Menéndez et al., 2020). Based on this approach, static, pseudo-  
551 dynamic and dynamic calculations have been performed. In the first and second cases, the  
552 obtained results coincide with analytical approaches. No analytical approach exists in the  
553 third case. Figure 19, illustrates the boulder toppling for the 3DEC model for a horizontal  
554 acceleration 0.105g (Muñiz-Menéndez et al., 2020).

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The use of numerical approaches, provided they are applied in a rigorous manner, can be quite useful to carry out particular analyses on the stability of boulders and group of boulders (Christianson et al., 1995). This could include consideration of some dynamic and coupled processes that difficult to address using standard approaches (Mendes et al., 2020, Lemos et al., 2011).

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## 5. Conclusions

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This paper is presented as an initial guideline for studies on granite boulder stability, or for other natural rock slope stability phenomena associated with irregular rock elements. Additionally, a number of issues still requiring further analysis are highlighted, with the aim of seeking improvement on our present capabilities to understand the actual instability behavior of granitic boulders.

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A combination of different remote sensing techniques (UAS photogrammetry and TLS) has been successfully demonstrated with the aim of developing accurate geometric models of boulders. These techniques are critical in providing a detailed geometrical representation of the rock element whose stability is at stake.

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The geometry and behavior of the contact zone of the boulder on the resting surface is another aspect that has shown to be important in the analysis of the stability of boulders against sliding or toppling. Sometimes it is not possible to fully constrain this geometry, so some assumptions regarding the geometry, roughness and strength of the contact are needed.

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External forces associated with water, ice-jacking or earthquakes can be considered the final trigger of the instability of some boulders, so suggestions are provided on how to compute the effects of these forces on the boulder stability. Moreover, the use of numerical models is briefly described, which can be helpful to manage this type of analysis.

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In summary, the authors have attempted to compile a number of relevant aspects playing a relevant role on the stability of granitic boulders in this document, with the aim of aiding the rock mechanics community in better assessing and predicting the mechanical stability of these natural structures.

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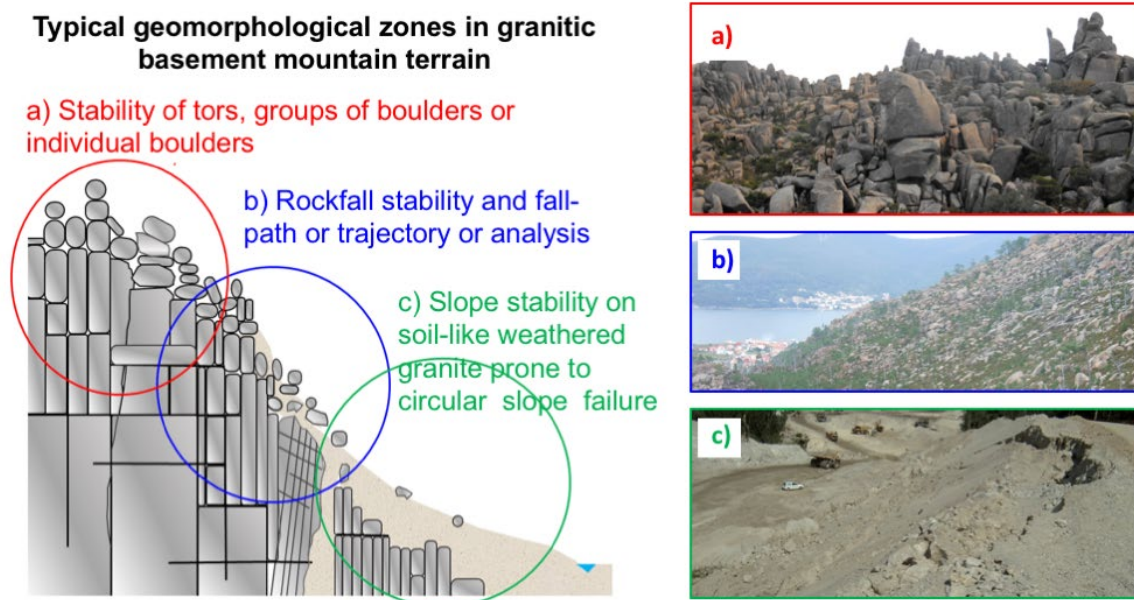
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## 726 FIGURES

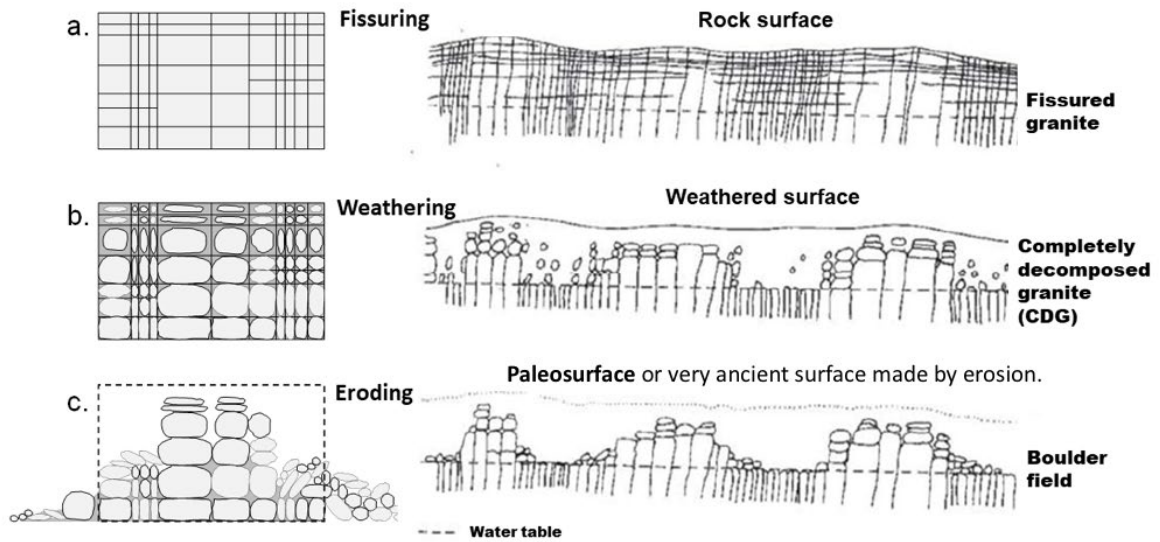
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729 Figure 1. Schematic diagram and accompanying pictures of different slope types in  
730 granitic terrain and associated potential geomechanical problems for each environment.

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## Process of granite boulder formation due to weathering



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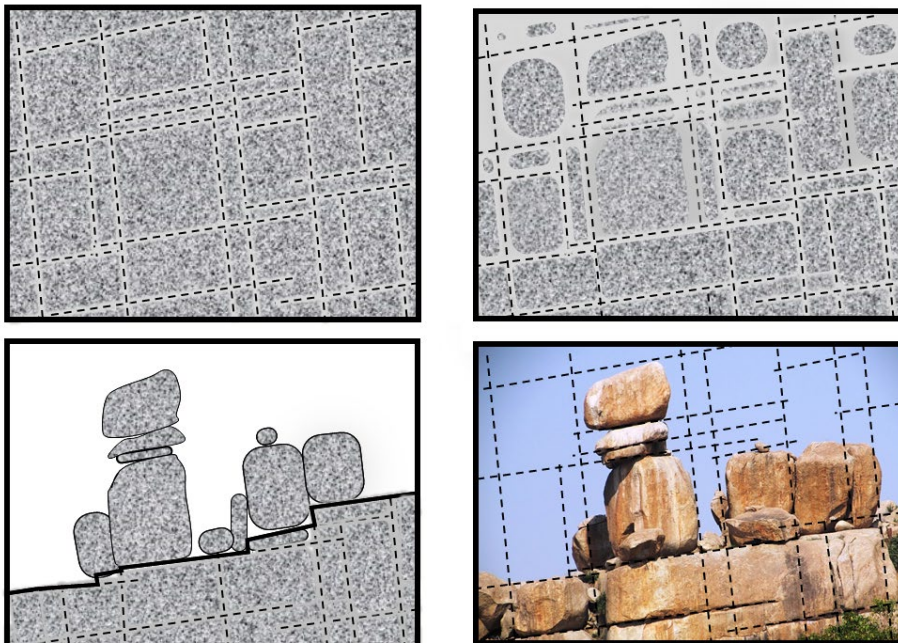
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Figure 2. Idealized sketches illustrating the evolution of granitic rock masses to produce boulder fields according to Linton (1955). a). Original granite orthogonally fissured rock mass; b). Spheroidal penetration of weathering; c). Ultimate stage with saprolite and clay removal, unveiling spheroidal weathering.



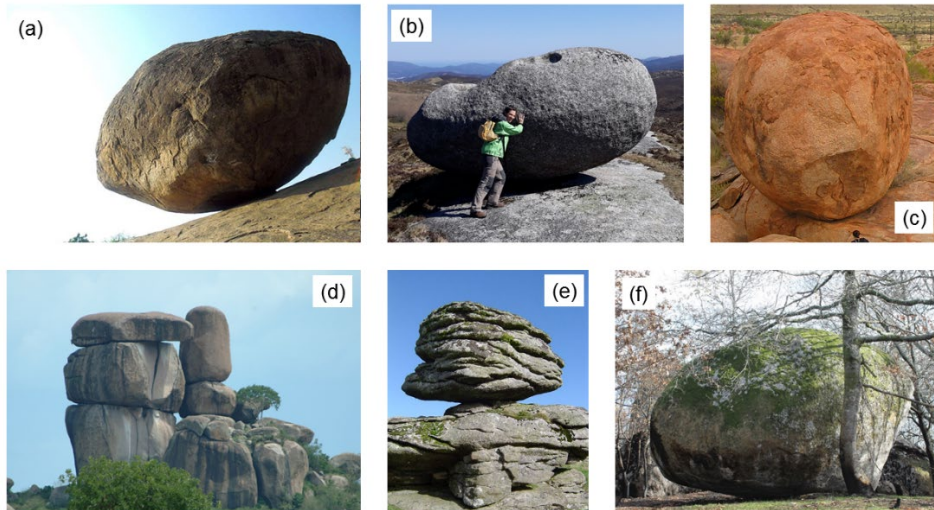
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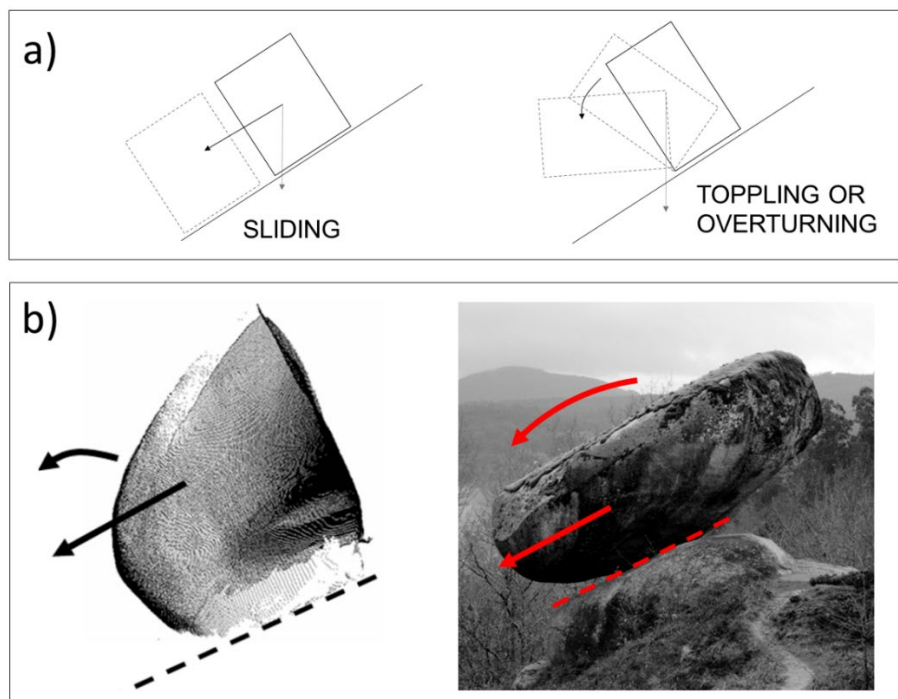
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Figure 3. Illustrative example of the tentative weathering and eroding process producing a group of granite boulders.



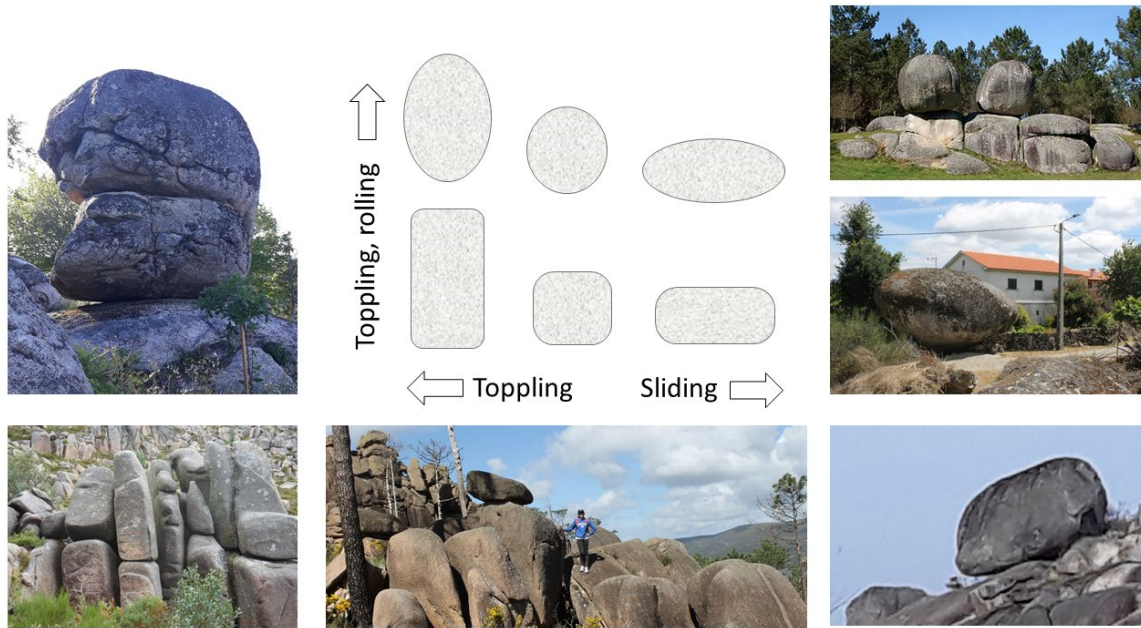
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Figure 4. Examples of granitic boulders in different parts of the world: a) The 250 t Krishna butter ball in the Kancheepuram district in India b) Kidney-shaped 100 t rocking stone in Abadín, Galicia (NW-Spain); c) Devil’s marble in the North Territories (Australia); d) Kit Mikayi or the stone of the first wife in Kisumu, Kenya, a 20 m high structure still attracting pilgrims, e) Logan stone (a rock which, through weathering, has become disjoined from the parent-rock and is pivoted upon it...) at Thornworthy Tor in UK; f) A 5 kt boulder in North Portugal. Source: photos b) and d) by the authors and a), c), d) and e) taken from (<https://commons.wikimedia.com>).



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Figure 5. a) Potential instability mechanisms of a boulder or a block lying on a tilted surface. b) Point cloud and picture images of two granite boulders lying on inclined planar surfaces or basal planes, whose stability was considered in previous studies (Alejano et al., 2010; Pérez-Rey et al., 2019a).



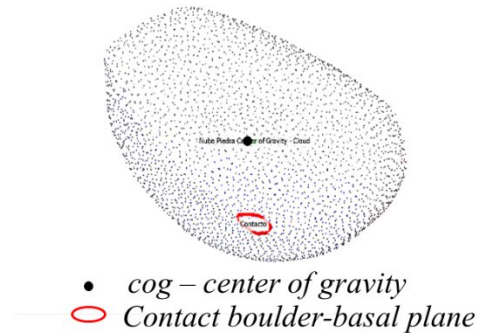
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Figure 6. Different geometry boulders. Slender blocks are more prone to topple, particularly if they show rounded corners. Rounded blocks tend to be more stable but may slide if they lay on basal planes more inclined than the contact friction angle.

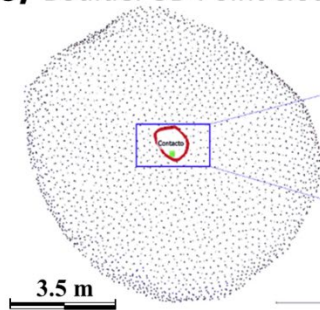
a) Boulder 3D-view



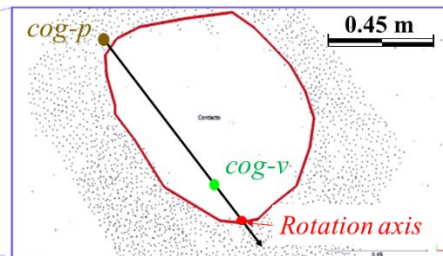
b) Boulder 3D Point cloud (3DPC)



c) Boulder 3D Point cloud top view



d) Enlarged top view of 3DPC

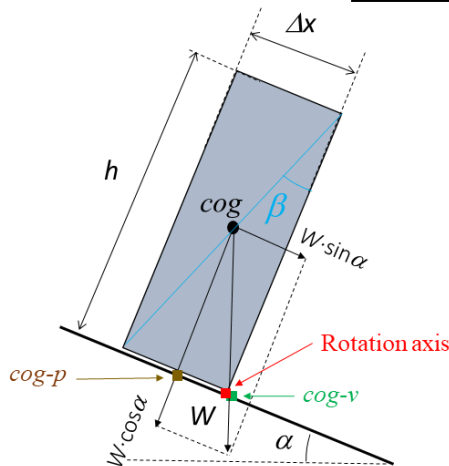


- cog-v – Vertical projection of the cog on the basal plane
- cog-p – Normal projection of the cog on the sliding plane

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Figure 7. a) 3D photogrammetric model and b) derived 3DPC of the ‘Pena do Equilibrio’ 350 t boulder in Spain, whose stability was studied by the authors. c & d) Plan view of the point cloud with area enlarged. Modified from Pérez-Rey et al., (2019a)

### Geometric regular block



$$FoS_{toppling} = \frac{\sum M_{stabilising}}{\sum M_{overturning}}$$

$$FoS_{toppling} = \frac{\frac{\Delta x}{2} W \cos \alpha}{\frac{h}{2} W \sin \alpha} = \frac{\Delta x}{h \tan \alpha} = \frac{\tan \beta}{\tan \alpha}$$

Slender blocks:  $\beta < \alpha$ , the block topples;  
 $FoS_{toppl} < 1$

Plump blocks:  $\beta > \alpha$ : the block is stable;  
 $FoS_{toppl} > 1$

*cog* – center of gravity

*cog-v* – Intersection of a vertical line passing through the cog with the sliding plane

*cog-p* – Normal projection of the cog on the sliding plane

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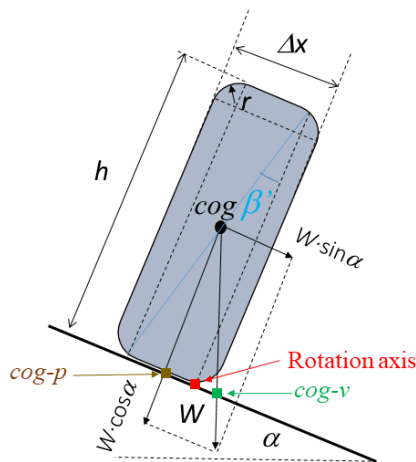
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Figure 8. Formulation of the stability of slab like block against toppling, based on the Goodman and Bray (1976) approach.

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### Geometric rounded cornered block



$$FoS_{toppling} = \frac{\sum M_{stabilising}}{\sum M_{overturning}}$$

$$FoS_{toppling} = \frac{(\frac{\Delta x}{2} - r) W \cos \alpha}{\frac{h}{2} W \sin \alpha} = \frac{\Delta x - 2r}{h \tan \alpha} = \frac{\tan \beta'}{\tan \alpha}$$

Very round corner:  $\beta' \ll \beta$  unstable

*cog* – center of gravity

*cog-v* – Intersection of a vertical line passing through the cog with the sliding plane

*cog-p* – Normal projection of the cog on the sliding plane

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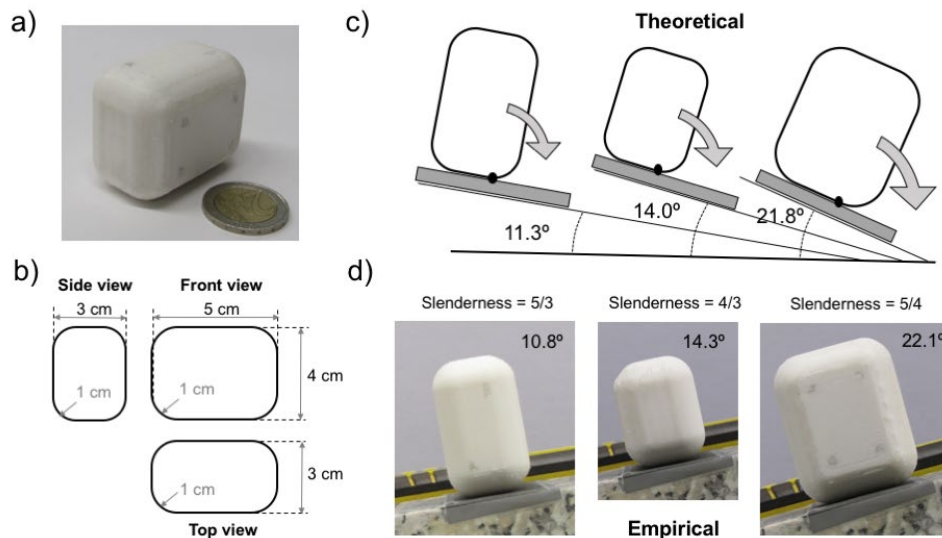
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Figure 9. Formulation of the stability analysis of slab like rounded corner block against toppling based on the Alejano et al. (2015) approach.

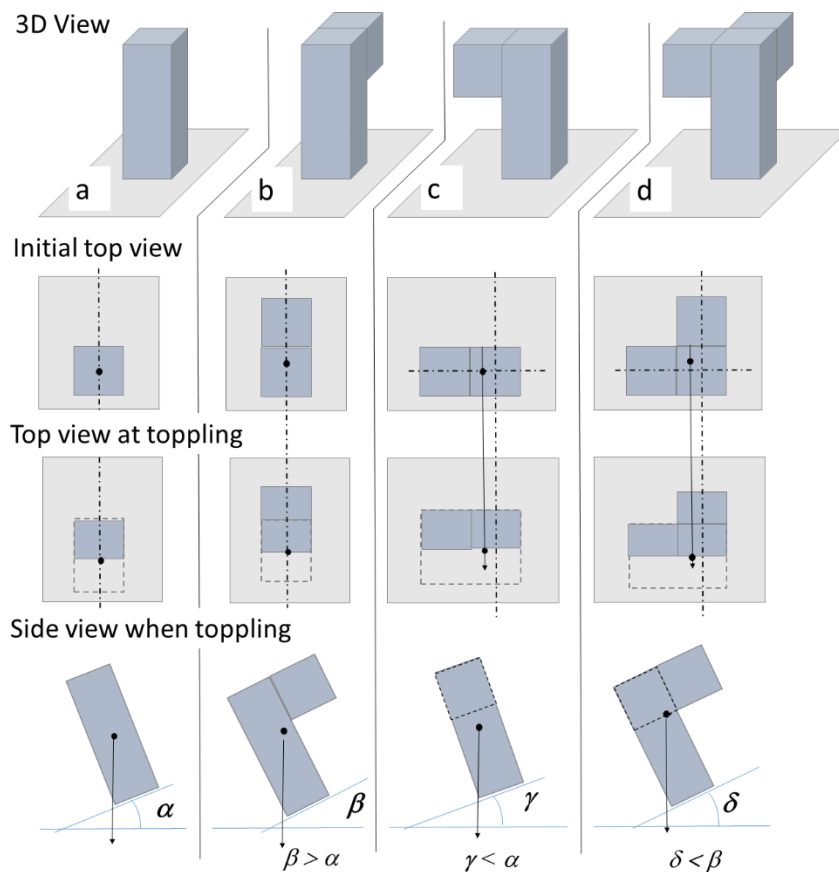
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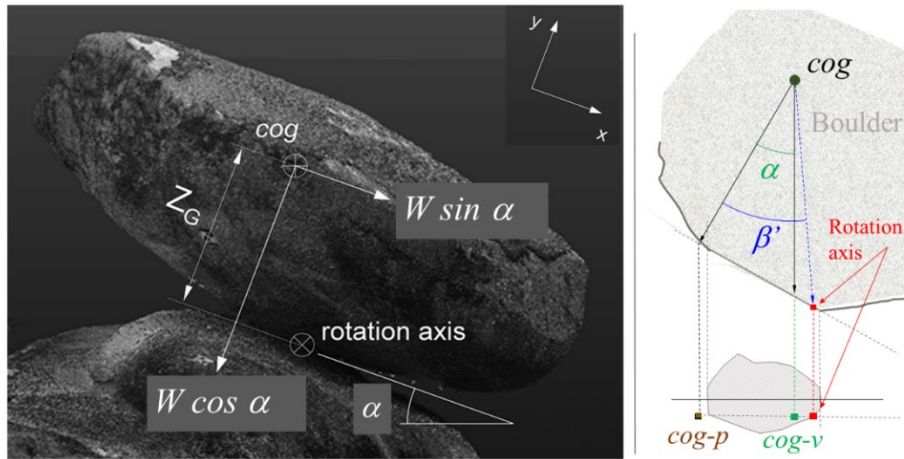
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Figure 10. Printed plastic element with fixed dimensions and perfectly rounded 1 cm radius corners and tilt tests showing theoretical and average empirical results.



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Figure 11. Image of various 3D elements *a*, *b*, *c* and *d* to be subjected to tilt testing to illustrate the role of geometry on toppling. On the upper row 3D view of the elements resting on a horizontal base to be tilted. On the second row, initial top view with the projection including the *cog*. On the third row, top view of the surface after tilting and in the moment of toppling and on the last row, side view of platform and element when toppling.



*cog* – Center of gravity

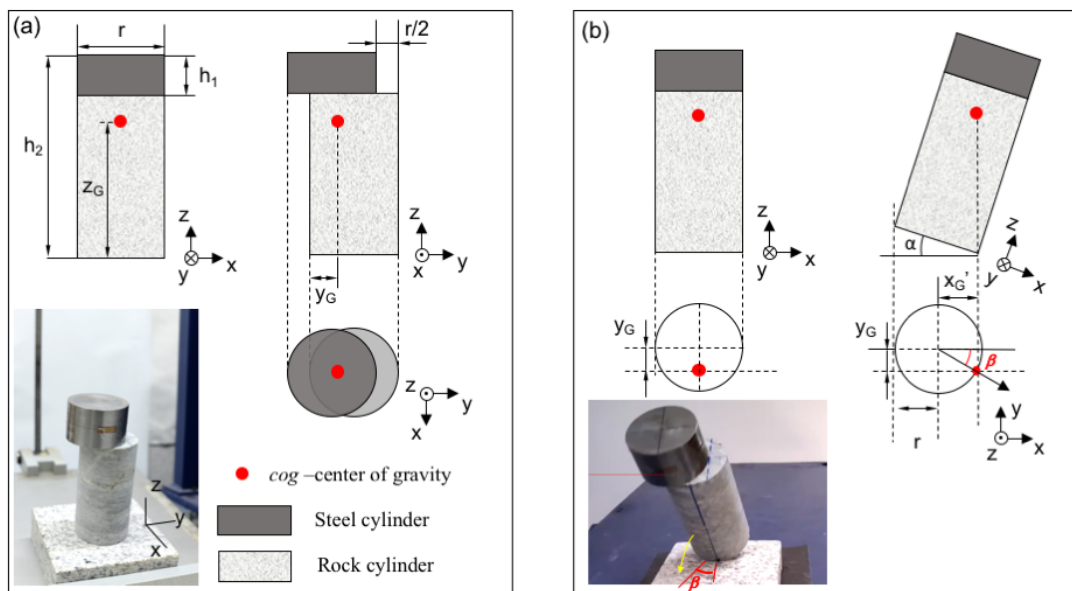
*cog-v* – Intersection of a vertical line passing through the *cog* with the sliding plane

*cog-p* – Normal projection of the *cog* on the sliding plane

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789 Figure 12. On the left hand side grayscale photograph, the location of the *cog* and forces  
 790 applied to the ‘Pena do Equilibrio’ boulder, whose stability was studied by the authors, is  
 791 shown. On the right hand side, a force diagram and projection of the contact area of this  
 792 boulder used to compute its stability against toppling are shown.

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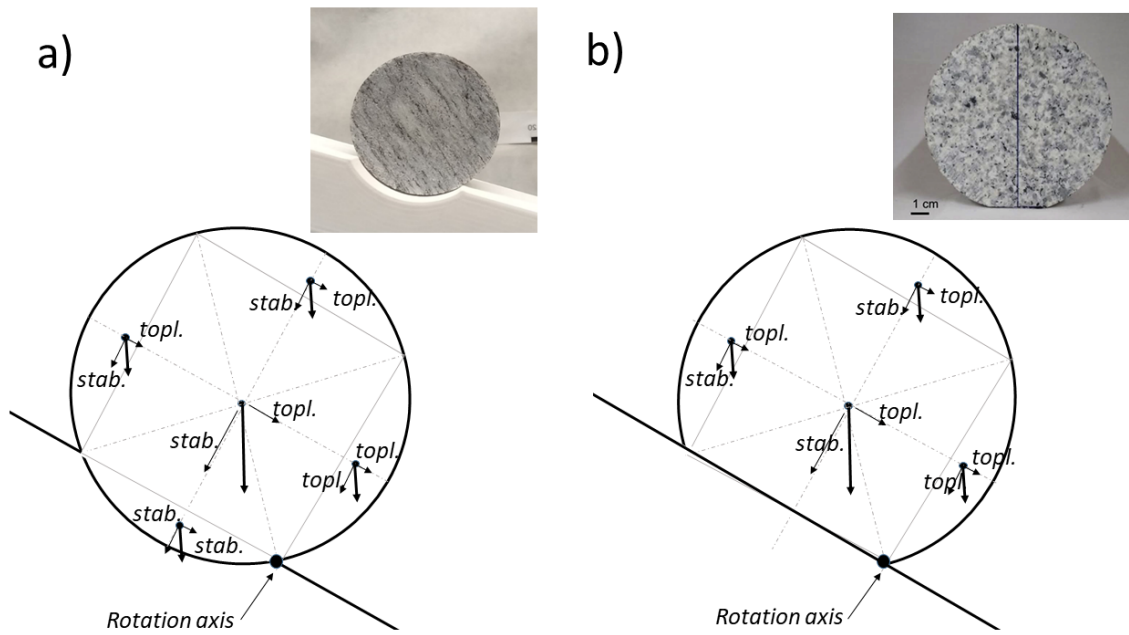


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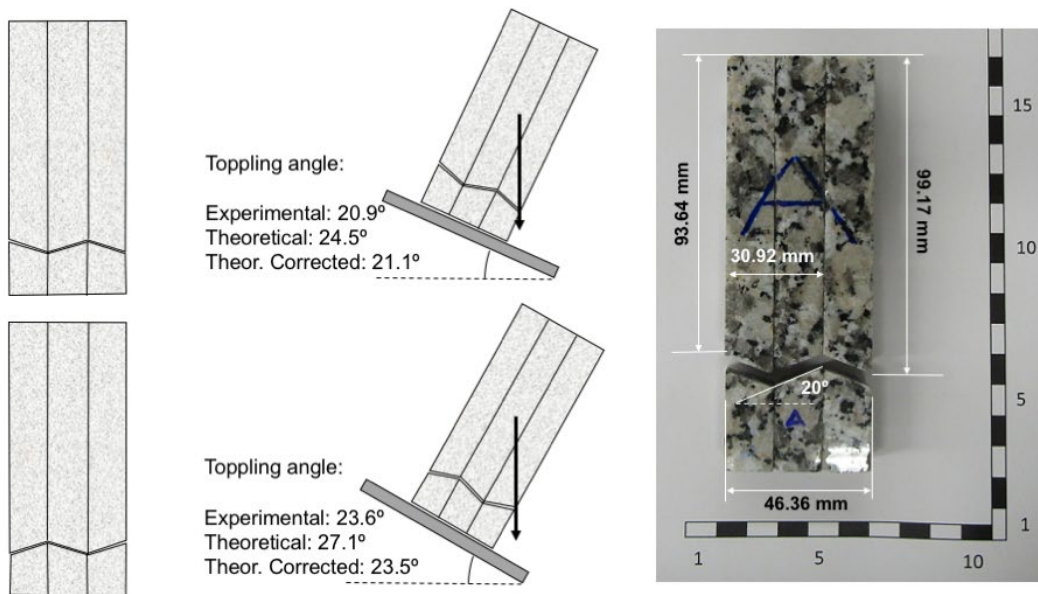


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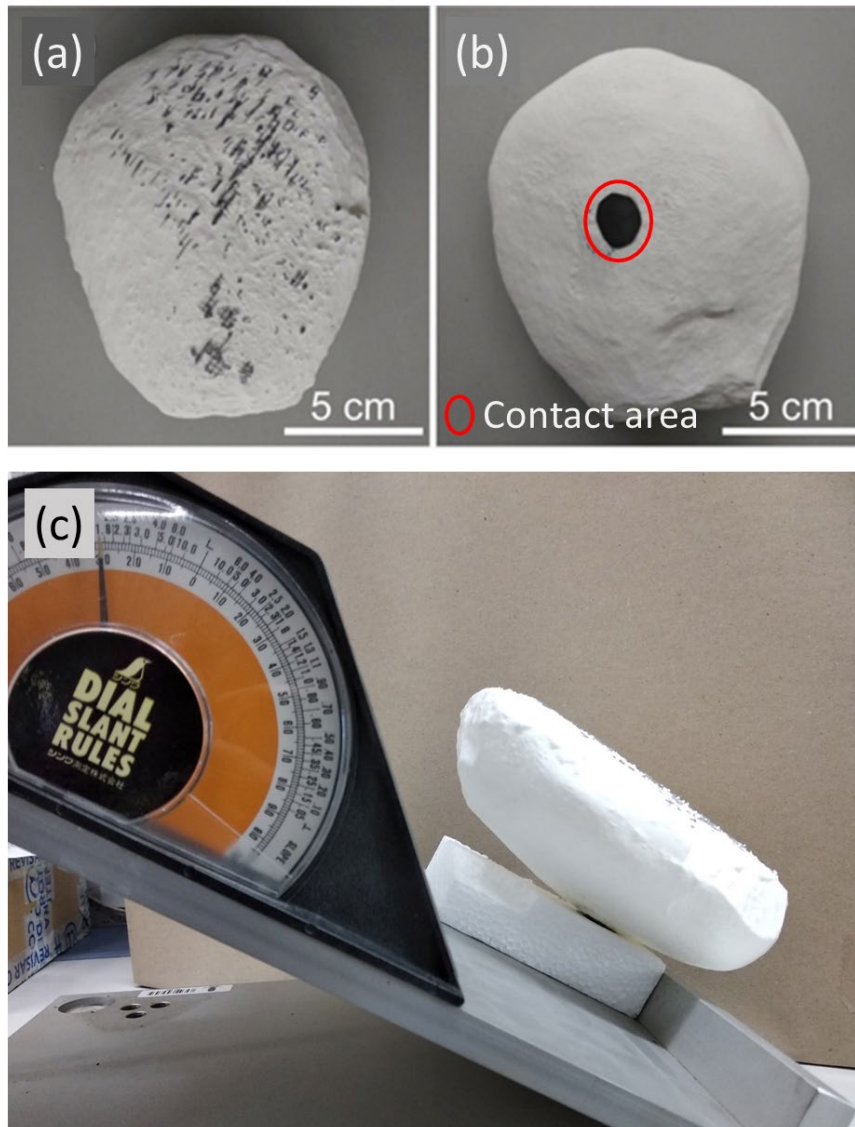
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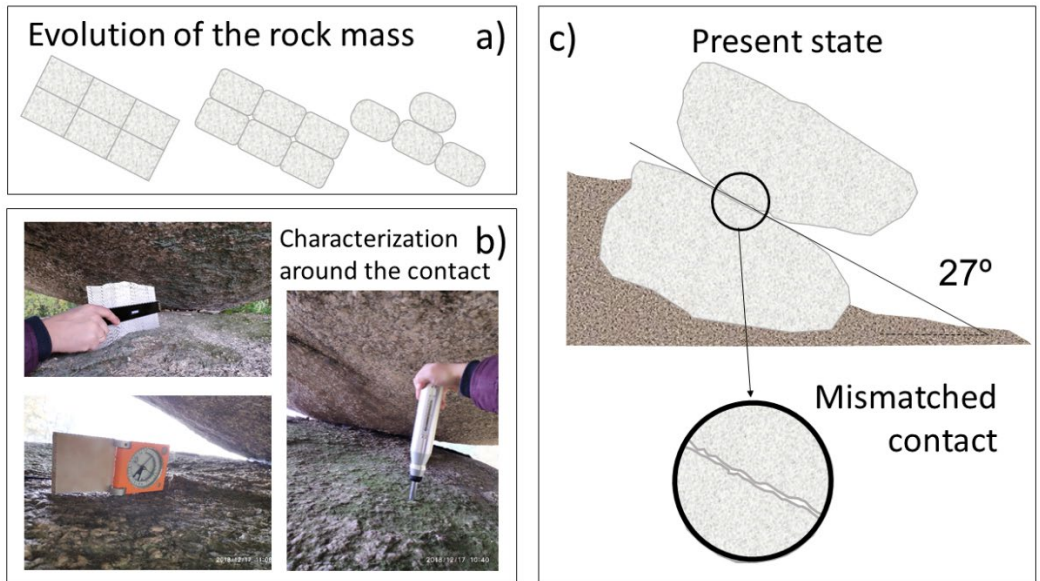
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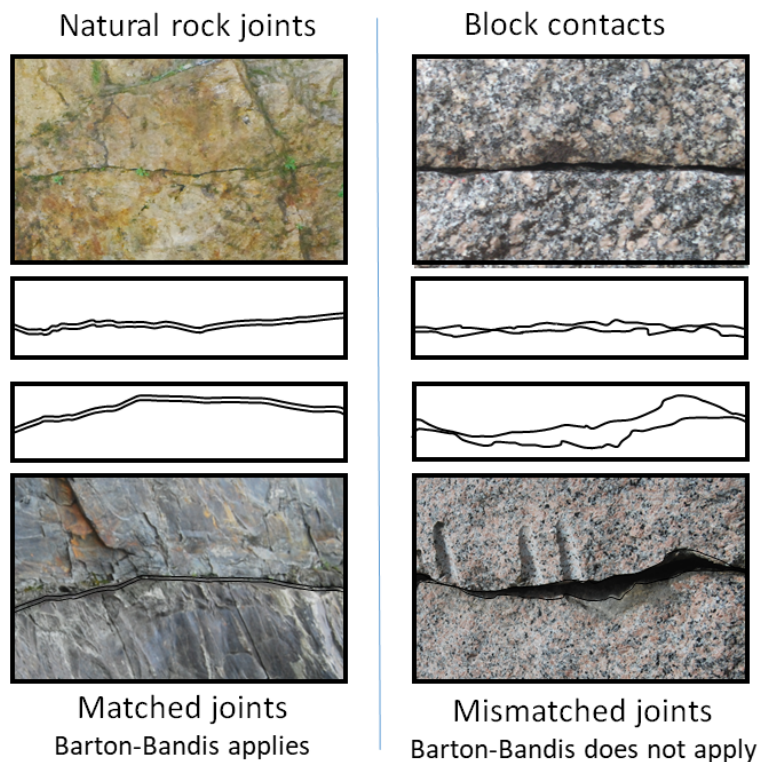
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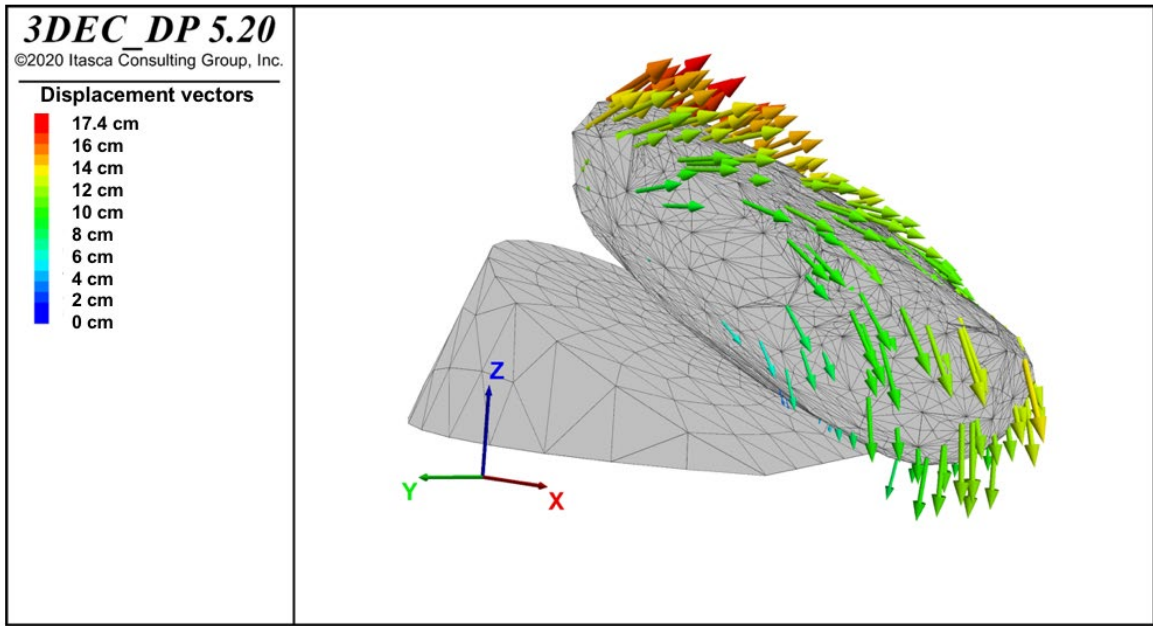
Figure 17. a) Sketch of the evolution of a granitic rock mass where starting from slab-like blocks, ellipsoidal boulders eventually occur. b) Detailed view of the small contact of a boulder with a basal plane and measurements taken (orientation, JRC and JCS). c) Sketch of the ‘Pena do Equilibrio’ boulder and detail illustrating the mismatched nature of the contact.



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Figure 18. Natural rock joints, which are matched joints on the left hand side and block contacts, which are mismatched joints, on the right hand side. The contact geometry is sketched for each case in the center of the Figure.

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Figure 19. Numerical result with code 3DEC of the toppling instability of the ‘Pena do Equilibrio’ boulder in a pseudo-dynamic analysis for the case of a large earthquake (horizontal acceleration 0.105g), illustrated by means of displacement vectors.

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### 837 TABLES

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Table 1. Results of toppling angles of a slab-like element with perfectly rounded corners. Tilt test and theoretical results.

Slenderness	Tilting results (°)	Average tests (°)	Theoretical result (°)
5/3	11, 12.3, 10.4, 9.5	10.8	11.3
4/3	15, 14.2, 13.4, 14.5	14.3	14.0
5/4	22.8, 22.2, 20.8, 22.7	22.1	21.8

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852 Figure 1. Schematic diagram and accompanying pictures of different slope types in granitic terrain.

853 Figure 2. Idealized sketches illustrating the evolution of granitic rock masses to produce boulder  
854 fields according to Linton (1955). a) Original granite orthogonally fissured rock mass; b)  
855 Spheroidal penetration of weathering; c) Ultimate stage with saprolite and clay removal, unveiling  
856 spheroidal weathering.

857 Figure 3. Illustrative example of the tentative weathering and eroding process producing a group of  
858 granite boulders.

859 Figure 4. Examples of granitic boulders in different parts of the world: a) The 250 t Krishna butter  
860 ball in the Kancheepuram district in India b) Kidney-shaped 100 t rocking stone in Abadín, Galicia  
861 (NW-Spain); c) Devil's marble in the North Territories (Australia); d) Kit Mikayi or the stone of  
862 the first wife in Kisumu, Kenya, a 20 m high structure still attracting pilgrims, e) Logan stone (a  
863 rock which, through weathering, has become disjoined from the parent-rock and is pivoted upon  
864 it...) at Thornworthy Tor in UK; f) A 5 kt boulder in North Portugal. Source: photos b) and d) by  
865 the authors and a), c), d) and e) taken from (<https://commons.wikimedia.com>).

866 Figure 5. a) Potential instability mechanisms of a boulder or a block lying on a tilted surface. b)  
867 Point cloud and picture images of two granite boulders lying on inclined planar surfaces or basal  
868 planes, whose stability was considered in previous studies (Alejano et al., 2010; Pérez-Rey et al.,  
869 2019a).

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884 base to be tilted. On the second row, initial top view with the projection including the *cog*. On the

885 third row, top view of the surface after tilting and in the moment of toppling and on the last row,  
886 side view of platform and element when toppling.

887 Figure 12. On the left hand side grayscale photograph, the location of the *cog* and forces applied to  
888 the ‘Pena do Equilibrio’ boulder, whose stability was studied by the authors, is shown. On the right  
889 hand side, a force diagram and projection of the contact area of this boulder used to compute its  
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